FOR OPSC & UPSC

GEOGRAPHY
OPTIONAL
PAPER 1& 2

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GEOGRAPHY

Paper-I

Principles of Geography

Section-A

Physical Geography

- i) Geomorphology: Origin of the earth, Physical conditions of the earth's interior; continental drift; isostasy; plate tectonics; mountain building; volcanism and earthquakes; weathering and erosion, Concepts of geomorphic cycles (Davis and Penck), Landforms associated with fluvial, arid, glacial, coastal and karst region, Polycyclic landforms.
- ii) **Climatology**: Temperature and pressure belts of the world; heat budget of the earth; atmospheric circulation; planetary and local winds; monsoons and jet streams; air masses and fronts; temperate and tropical cyclones; types and distribution of precipitation; Koppen's and Thornthwaite's classification of world climate; hydrological cycle; climatic change.
- iii) Oceanography: Bottom topography of the Atlantic, Indian and Pacific Oceans; temperature and salinity of the oceans; ocean deposits; ocean currents and tides; marine resources and their utilizations, Coral reefs;
- iv) **Biogeography**: Genesis of soils; classification and distribution of soils; soil profile; soil erosion and conservation; factors influencing world distribution of plants and animals; problems of deforestation and conservation measures; social forestry, agroforestry.
- v) Environmental Geography: Concept and types of environment, Environmental degradation and management. Ecosystems and their management; Energy flow and Biogeo- chemical cycles, Global ecological imbalances—problems of pollution, global warming, reduction in bio-diversity and depletion of forests.

Section-B

Human Geography

- i) Perspectives in Human Geography: Areal differentiation; regional synthesis; dichotomy and dualism; environmentalism; quantitative revolution and locational analysis; radical, behavioural, human and welfare approaches; Cultural regions of the world, Human development indicators. ii) Economic Geography: World economic development—measurement and problems; world resources and their distribution; energy crisis; the limits to growth; World agriculture—typology of agricultural regions; Von-Thunen's theory of agricultural location; World industries—locational patterns and locational theories of Weber; Hoover, Losch and Smith, Patterns of world trade. iii) Population Geography: Growth and distribution of world population; demographic attributes; causes and consequences of migration; concepts of over—, under— a n d optimum population; world population problems. Races of man kind.
- **iv)** Settlement Geography: Types and patterns of rural settlements; hierarchy of urban settlements; Cristaller's Central Place Theory, concept of primate city and rank-size rule; functional classification of towns; sphere of urban influence; rural-urban fringe; satellite town; problems of urbanisation.
- v) Regional Planning: Concept of region; types of regions and methods of regionalisation; growth centres and growth poles; regional imbalances; multi-level planning; planning for sustainable development. Rostov Model of Stages of Growth.

Note: Candidates will be required to answer one compulsory map question pertinent to subjects covered by this paper.



Paper-II

Geography of India with special reference to Odisha Section-A

- i) Physical Aspects: Structure and relief; drainage system and watersheds; physiographic regions; mechanism of Indian monsoons; tropical cyclones and western distrubances; floods and droughts; climatic regions; natural vegetation, soil types and their distributions.
- ii) Resources: Concept and types of resources, land, water, energy, minerals, and biotic resources, their distribution, utilisation and conservation; energy crisis.
- iii) Agriculture: Infrastructure—irrigation, seeds, fertilizers, power; Types of crops agricultural productivity, agricultural intensity, crop combination, land capability; agroand social forestry; Green Revolution its socio-economic and ecological implications; significance of dry farming; livestock resources and White Revolution; Blue Revolution; agricultural regionalisation; agroclimatic zones.
- iv) Industry: History of industrial development; locational factors of cotton, jute, iron and steel, fertilizer and paper, industries, industrial complexes and industrial regionalisation; new industrial policy; role of multinationals, liberalization and globalisation.
- v) Transport, Communication and Trade: Road, railway, waterway, airway and pipeline networks and their complementary roles in regional development; growing importance of ports on national and foreign trade, trade balance; free trade and export promotion zones; developments in communication technology and its impact on economy and society.

Section-B

- i) Cultural Setting: Racial and ethnic diversities; major tribes, tribal areas and their problems; role of language, religion and tradition in the formation of cultural regions; growth, distribution and density of population; demographic attributes—sex-ratio, age structure, literacy rate, work-force, dependency ratio and longivity; migration (interregional, intra-regional and international) and associated problems, population problems and policies.
- ii) **Settlements**: Types, patterns and morphology of rural settlements; urban development; Census definition of urban areas; morphology of Indian cities; functional classification of Indian cities; conurbations and metropolitan regions; urban sprawl; slums and associated problems; town planning; problems of urbanisaiton.
- iii) Regional Development and Planning: Experience of regional planning in India; Five Year Plans; integrated rural development programmes; Panchayati Raj and decentralised planning; command area development; watershed management; planning for backward area, desert, drought-prone, hill and tribal area development; multi-level planning; geography and regional planning.
- iv) Political Aspects: Geographical basis of Indian federalism; state reorganisation; regional consciousness and national integration; international boundary of India and related issues; disputes on sharing of water resources; India and geopolitics of the Indian Ocean.
- v) Contemporary Issues: Environmental hazards—landslides, earthquakes, Tsunami, cyclones, floods and droughts, epidemics. Issues related to environmental pollution; changes in patterns of land use; principles of environmental impact assessment and environmental management; population explosion and food security; environmental degradation; Disasters in India and their management. Problems of agrarian and industrial unrest; regional disparities in economic development; concept of sustainable growth and development.

Note: Candidates will be required to answer one compulsory map question pertinent to subjects covered by this paper.



PHYSICAL GEOGRAPHY

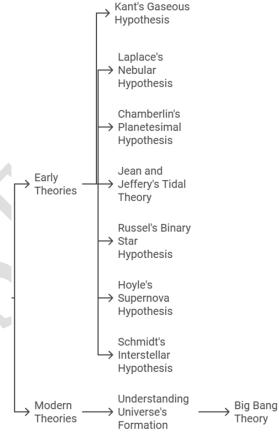
GEOMORPHOLOGY

ORIGIN OF THE EARTH

The origin of the Earth is a complex topic that has intrigued scientists and philosophers for centuries. Various theories have emerged to explain how our planet formed approximately 4.6

billion years ago. Theories for the earth's origin has been classified into two categories i.e., early theories and modern theory which are as follows:

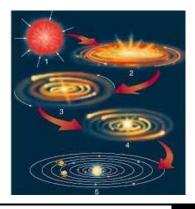
- Early Theories: Early Theories explained how the earth was formed. The early theories were the Gaseous hypothesis of Kant, the Nebular Hypothesis of Laplace, the Planetesimal hypothesis of Chamberlin, Jean and Jeffery's tidal theory/Gravitational Theory, Russel's binary star hypothesis, Hoyle's supernova hypothesis, Schmidt's interstellar hypothesis.
- Modern Theories: Modern theories state the need to find how the universe has formed then automatically we will be able to find the formation of the earth. Early theories focused solely on the evolution of the earth and planets, whereas recent theories attempt to solve the questions of the universe's creation. Big Bang theory is an example of modern theory.



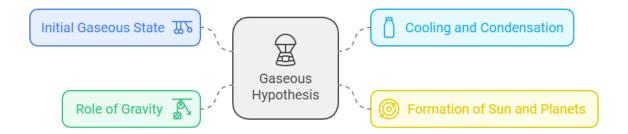
1. Theories of Origin: The theories regarding the origin of the Earth can be broadly categorized into early hypotheses and modern scientific theories.

A. Early Theories

Gaseous Hypothesis (Kant): Proposed by Immanuel Kant in the 18th century, this theory posits that the Earth formed from a primordial cloud of gas and dust. Kant suggested that this cloud slowly rotated and condensed due to gravitational forces, leading to the formation of the Sun and planets. According to Kant, the Earth was formed from small, cold particles that aggregated over time.







Assumptions and Postulates

- 1. **Primordial Matter**: Kant assumed that the universe was filled with primordial matter, consisting of small, hard, and cold particles.
- 2. **Gravitational Attraction**: These particles were drawn together by gravitational attraction, leading to collisions that generated heat.
- 3. Formation of a Nebula: The heat and collisions caused the primordial matter to transform into a vast, hot, rotating nebula.
- 4. **Centrifugal Force**: As the nebula rotated, centrifugal force caused rings of material to be thrown off, which eventually condensed to form planets.
- 5. **Solar System Formation:** The remaining central mass became the Sun, and the planets formed from the rings of material.

Criticisms

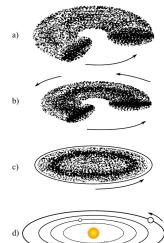
- Origin of Primordial Matter: Kant's hypothesis did not explain the origin of the primordial matter itself.
- 2. Lack of Empirical Evidence: At the time, there was no empirical evidence to support the hypothesis.
- 3. **Simplistic Model**: The model was considered too simplistic and did not account for the complexities observed in the solar system.

Present Relevance

- 1. Foundation for Future Theories: Kant's hypothesis laid the groundwork for future theories, such as Laplace's Nebular Hypothesis.
- 2. Scientific Approach: It was one of the first theories to apply Newton's laws of gravity and motion to explain the formation of the solar system.
- 3. **Historical Significance**: While not entirely accurate, Kant's hypothesis is historically significant as it marked a shift towards scientific explanations for cosmic phenomena

Nebular Hypothesis (Laplace): The nebular hypothesis, also known as the nebular theory, is a scientific theory that explains how the solar system formed. It was developed by French astronomer and mathematician Pierre-Simon Laplace in 1796.

The theory suggests that the solar system formed from a giant cloud of gas and dust, called a nebula, that collapsed under its own gravity.





As the nebula collapsed, it spun faster and flattened into a disk shape. The material in the center became hotter and denser, eventually forming the sun. The rest of the disk formed the planets and moons.

Nebular	Rotating	Gravitational	Protostar	Protoplanet	、 Planet
Hypothesis (Nebula	Collapse	Formation	Creation	Evolution

Assumptions and Postulates

- 1. **Primordial Nebula**: The solar system originated from a large, rotating cloud of gas and dust, known as a nebula.
- 2. **Gravitational Collapse**: The nebula began to collapse under its own gravity, causing it to spin faster and flatten into a disk.
- 3. Formation of the Sun: The central region of the collapsing nebula became increasingly hot and dense, eventually forming the Sun.
- 4. Planetary Formation: The remaining material in the disk began to coalesce into planetesimals, which collided and stuck together to form planets.
- 5. **Conservation of Angular Momentum**: The rotation of the nebula ensured that the planets formed in the same plane and direction as the original nebula's rotation.

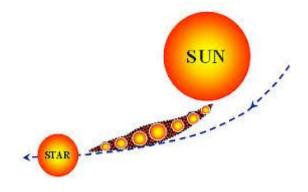
Criticisms

- 1. Angular Momentum Distribution: The hypothesis struggled to explain the distribution of angular momentum between the Sun and the planets.
- 2. Formation of Planetesimals: The process by which small particles coalesced into larger planetesimals was not well understood.
- 3. **Temperature Variations**: The hypothesis did not account for the temperature variations needed to form different types of planets (rocky and gaseous).

Present Relevance

- 1. Foundation for Modern Theories: The Nebular Hypothesis laid the groundwork for modern theories of planetary formation, such as the Solar Nebula Disk Model (SNDM).
- 2. Scientific Approach: It introduced the idea of a natural, physical process for the formation of the solar system, moving away from mythological explanations.
- 3. Observational Evidence: Observations of protoplanetary disks around young stars provide evidence supporting the basic concepts of the Nebular Hypothesis.

Planetesimal Hypothesis: The Planetesimal Hypothesis, proposed by Thomas Chamberlin and Forest Moulton in the early 20th century, is an intriguing theory about the formation of the solar system This hypothesis suggests that planets formed from small bodies called planetesimals. These planetesimals coalesced through gravitational attraction over time, eventually forming larger planetary bodies. This idea emphasizes the role of collisions and accretion in planet formation.





Assumptions and Postulates

- 1. Near-Collision Event: The hypothesis suggests that a passing star came close to the Sun, causing tidal forces to pull material from both stars.
- 2. Formation of Planetesimals: The material pulled from the stars condensed into small, solid bodies called planetesimals.
- 3. Accretion Process: These planetesimals collided and stuck together over time, gradually forming larger bodies, including planets and moons.
- 4. **Gravitational Influence**: The gravitational forces of the Sun and the passing star played a crucial role in shaping the orbits and distribution of the planetesimals.

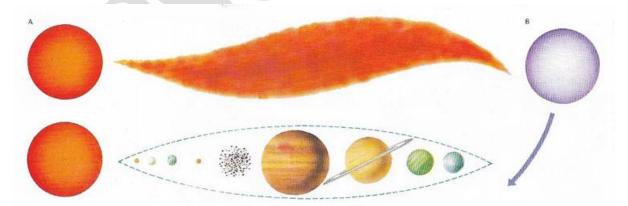
Criticisms

- 1. Improbability of Stellar Encounters: The hypothesis relies on a close encounter between stars, which is statistically very unlikely given the vast distances between stars.
- 2. Expansion of Hot Gases: Critics argue that the hot gases pulled from the stars would expand and dissipate rather than condense into solid bodies.
- 3. Angular Momentum Distribution: The hypothesis does not adequately explain the distribution of angular momentum in the solar system.
- 4. Formation of Only Eight Planets: It fails to explain why only eight planets formed and why they have such different sizes and compositions.

Present Relevance

- 1. Foundation for Modern Theories: Despite its flaws, the Planetesimal Hypothesis laid the groundwork for modern theories of planetary formation, such as the Solar Nebula Theory.
- 2. **Concept of Planetesimals**: The idea of planetesimals as building blocks of planets remains a key concept in current models of planetary formation.
- 3. **Understanding Accretion**: The hypothesis contributed to our understanding of the accretion process, where small particles collide and stick together to form larger bodies

Tidal Theories of Jeans and Jeffrey: The Tidal Theories of Jeans and Jeffrey, proposed by Sir James Jeans in 1919 and later modified by Harold Jeffreys in 1929, offer an intriguing explanation for the formation of the solar system. Here's a detailed look at their assumptions, postulates, criticisms, and present relevance:



Assumptions and Postulates

- 1. Close Stellar Encounter: The theory suggests that the solar system was formed from the Sun and another massive star, referred to as the "intruding star."
- 2. **Gravitational Interaction**: As the intruding star passed close to the Sun, its gravitational pull created massive tidal forces on the Sun's surface.



- 3. **Ejection of Material**: These tidal forces caused a large amount of material to be ejected from the Sun in the form of a cigar-shaped filament.
- 4. Formation of Planets: The ejected material cooled and condensed to form planets. The central, thicker part of the filament formed the larger planets like Jupiter and Saturn, while the thinner ends formed the smaller planets.

Criticisms

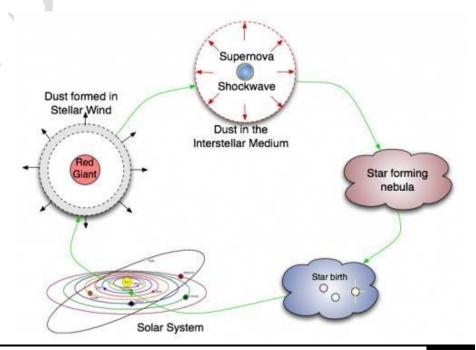
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Present Relevance

- 1. **Historical Significance**: The Tidal Theories of Jeans and Jeffrey are historically significant as they represent an early attempt to explain the formation of the solar system using gravitational interactions.
- 2. **Foundation for Modern Theories**: Despite its flaws, the theory laid the groundwork for future research and the development of more accurate models of planetary formation.
- 3. Concept of Tidal Forces: The idea of tidal forces influencing the formation of celestial bodies remains relevant in modern astrophysics

Interstellar Dust Hypothesis of Otto Schmidt

The Interstellar Dust Hypothesis, proposed by Russian scientist Otto Schmidt in 1943, offers a compelling explanation for the origin of the solar system and its characteristics. Schmidt proposed that the Sun passed through a dense interstellar cloud and became enveloped in a cloud of dust and gas. The dust particles condensed and organized, eventually forming the Earth and other planets.





This hypothesis addresses several complex problems regarding how the Earth and other celestial bodies formed from primordial matter in the universe.

Key Concepts of the Interstellar Dust Hypothesis

- Dark Matter and Dust Clouds: Schmidt suggested that the universe contains vast amounts
 of dark matter in the form of gas and dust clouds, which are remnants from stars and
 meteors. These clouds, often referred to as "interstellar dust," play a crucial role in the
 formation of celestial bodies.
- 2. Galactic Revolution: According to Schmidt, as the Sun moved through the galaxy, it captured this dark matter. The gravitational pull of the Sun attracted these gas and dust particles, which began to revolve around it due to their own angular momentum.
- 3. Formation of a Disc: The captured interstellar dust and gas particles coalesced into a flat disc shape around the Sun. This disc was characterized by three types of motions:
 - **Gravitational Force**: The Sun's gravity exerted a force on the disc, pulling particles toward it.
 - Rotational Motion: The Sun's own rotational motion contributed to the dynamics of the disc
 - Angular Momentum: The inherent angular momentum of the dark matter influenced how these particles interacted.
- 4. Planet Formation: Over time, as these dust particles collided and merged, they formed larger bodies that eventually became planets. This process involved complex interactions among gravitational forces, rotational dynamics, and the properties of the materials involved.

Evaluation of Schmidt's Hypothesis:

The Interstellar Dust Hypothesis addresses many peculiar characteristics observed in our solar system:

- Orbital Patterns: It explains why planets have near-circular orbits that lie within a similar plane (the ecliptic plane), as they formed from a rotating disc.
- Size Distribution: The hypothesis accounts for the placement of planets according to their size and density, suggesting that different materials coalesced based on their proximity to the Sun and their physical properties.
- Angular Momentum Distribution: It provides insights into why there is a significant distribution of angular momentum among planets, consistent with their formation from a rotating disc.

B. Modern Theories



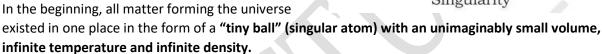
BIG BANG THEORY

Scientists in early 20th Century took up the problems of origin of universe rather than that of just the earth or the planets. In 1927, a Belgian cosmologist Georges Lemaitre proposed the

Big Bang theory. At some point in the past, he argued, the entire mass of the universe would have been concentrated into a single point from which the very fabric of space and time originated.

 The Big Bang hypothesis states that all of the current and past matter in the Universe came into existence at the same time, roughly 13.8 billion years ago. At this time, all matter was compacted into a very small ball with infinite density and intense heat called a Singularity.

The Big Bang Theory considers the following stages in the development of the universe.



At the Big Bang the "tiny ball" exploded violently. This led to a huge expansion. It is now generally accepted that the event of big bang took place 13.7 billion years before the present. The expansion continues even to the present day. As it grew, some energy was converted into matter. There was particularly rapid expansion within fractions of a second after the bang. Thereafter, the expansion has slowed down. Within first three minutes from the Big Bang event, the first atom began to form.

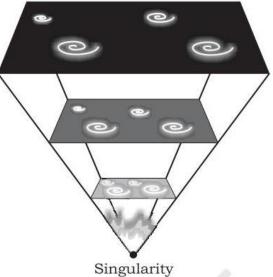
Within 300,000 years from the Big Bang, temperature dropped to 4,500K and gave rise to atomic matter. The universe became transparent.

The expansion of universe means increase in space between the galaxies. An alternative to this was Hoyle's concept of steady state. It considered the universe to be roughly the same at any point of time. However, with greater evidence becoming available about the expanding universe, scientific community at present favours argument of expanding universe.

Evidences of Big Bang Theory

The Big Bang Theory is supported by several key pieces of evidence that have been observed and studied over the years. Here are some of the most significant ones:

- Cosmic Microwave Background Radiation (CMBR): This is the afterglow of the Big Bang, a
 faint radiation that fills the universe and can be detected in every direction
- 2. Redshift of Galaxies: This redshift of galaxies indicates that the universe is expanding, which supports the idea that it started from a single point.
- 3. Abundance of Light Elements: The Big Bang Theory predicts the proportions of light elements (hydrogen, helium, and lithium) that should have been produced in the first few minutes after the Big Bang.





- 4. Large-Scale Structure of the Universe: The distribution of galaxies and galaxy clusters throughout the universe can be explained by the initial density fluctuations in the early universe, as predicted by the Big Bang Theory.
- 5. **Hubble's Law**: This law states that the velocity at which a galaxy is receding from us is directly proportional to its distance from us.

Contradictions and challenges:

- In the 1920s and 1930s almost every major cosmologist **preferred an endless steady state** universe instead of expanding universe.
- Singularity Problem: The theory begins with a singularity, a point of infinite density and temperature where the laws of physics as we know them break down. This concept is difficult to reconcile with our current understanding of physics.
- Flatness Problem: The observed flatness of the universe requires fine-tuning of initial conditions, which is not naturally explained by the Big Bang Theory.
- Dark Matter and Dark Energy: The Big Bang Theory does not explain the nature of dark matter and dark energy, which together constitute about 95% of the universe's mass-energy content.
- Initial Conditions: The theory does not explain why the initial conditions of the universe were such that they led to the formation of galaxies, stars, and planets. It assumes these conditions without providing a mechanism for their origin.
- Too much energy: Most of the universe is empty space: a vacuum which is defined as a
 volume containing no particles, force fields, nor waves. By definition a vacuum has no
 energy. However, the Big Bang theory requires both in its early phases and in its later phases
 that the vacuum must have some energy. The law of conservation of energy demands that
 energy cannot be created nor destroyed.
- Too complex, too early: The universe has too many large structures, to be created in 10-20 billion years. We know the rate of expansion; thus, we can get a rough estimate on how long it would take for them to form. In order for these to form, it would take about 100 billion years.
 - 1. Discuss the various hypotheses and theories regarding the origin of the Earth.
 - 2. Critically analyze the Planetesimal Hypothesis and its relevance in modern astronomy.
 - 3. Evaluate the Tidal Theories of Jeans and Jeffrey in the context of the origin of the Earth.
 - 4. Discuss the role of volcanic outgassing in the formation of Earth's early atmosphere.
 - 5. Analyze the contributions of different scientists to the understanding of the Earth's origin.
 - 6. Discuss the Inter-stellar Dust Hypothesis of Otto Schmidt and its implications for the formation of the Earth.
 - 7. Explain the Nova Star Hypothesis of Hoyle and Lyttleton and its relevance to the origin of the Earth.
 - 8. Explain the Big Bang Theory and discuss the evidence supporting it.
 - 9. Evaluate the significance of redshift in understanding the expanding universe as proposed by the Big Bang Theory.
 - 10. Discuss the contributions of key scientists to the development of the Big Bang Theory.



FORMATION OF PLANETS

The following are considered to be the stages in the development of planets:

- The stars are localised lumps of gas within a nebula. The gravitational force within the lumps leads to the formation of a core to the gas cloud and a huge rotating disc of gas and dust develops around the gas core.
- In the next stage, the gas cloud starts getting condensed and the matter around the core develops into small rounded objects.
 These small-rounded objects by the process of cohesion develop into what is called planetesimals. Larger bodies start forming by collision, and gravitational attraction causes the material to stick together. Planetesimals are a large number of smaller bodies.
- 3. In the final stage, these large number of small planetesimals accrete to form a fewer large bodies in the form of planets.

The planet earth initially was a barren, rocky and hot object with a thin atmosphere of hydrogen and helium. This is far from the present day picture of the earth.

Formation of Earth's Atmosphere: The formation of Earth's atmosphere is a fascinating and complex process that took place over billions of years. Here's an overview of how it happened:

Early Atmosphere

1. Primordial Atmosphere: Shortly after Earth formed about 4.5 billion years ago, its initial atmosphere was likely composed of hydrogen and helium, the most abundant elements in the universe. However, this atmosphere was quickly lost to space due to the Sun's intense solar wind and Earth's weak gravity at the time.

Primordial Atmosphere Lost to Space Volcanic Outgassing Formation of Oceans Photosynthesis by Cyanobacteria Ozone Laver Formation Biological Geological Processes Present Atmosphere

Secondary Atmosphere

- 2. Volcanic Outgassing: As the Earth cooled and solidified, volcanic activity became intense. Volcanic eruptions released gases trapped in the planet's interior, including water vapor, carbon dioxide, nitrogen, methane, ammonia, and other trace gases. This process, known as outgassing, created a new, thicker atmosphere.
- 3. Formation of Oceans: The water vapor released by volcanic activity eventually condensed to form clouds and precipitated as rain, leading to the formation of Earth's oceans. This process also helped to remove some carbon dioxide from the atmosphere, as it dissolved in the water and formed carbonate rocks.

Development of the Modern Atmosphere

4. Photosynthesis and Oxygenation: Around 2.5 billion years ago, cyanobacteria (blue-green algae) began to perform photosynthesis, a process that converts carbon dioxide and water into oxygen and glucose using sunlight. This led to the gradual accumulation of oxygen in the atmosphere, known as the Great Oxidation Event.



- 5. Ozone Layer Formation: As oxygen levels increased, some of it was converted into ozone (O3) in the upper atmosphere. The ozone layer formed and began to protect the Earth from harmful ultraviolet (UV) radiation, allowing more complex life forms to evolve.
- Biological and Geological Processes: Over time, the balance of gases in the atmosphere has been influenced by various biological and geological processes, including the carbon cycle, respiration, and the burial of organic matter.

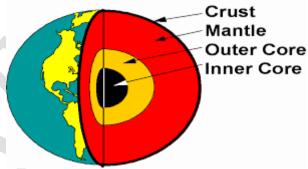
Present Atmosphere

Today, Earth's atmosphere is composed of approximately 78% nitrogen, 21% oxygen, and trace amounts of other gases such as argon, carbon dioxide, neon, and methane. This composition supports a diverse range of life forms and helps regulate the planet's climate.

PHYSICAL CONDITIONS OF THE EARTH'S INTERIOR

The Earth's interior is a complex and dynamic system with varying physical conditions as you move from the surface to the core. Here's a brief overview:

1. Crust: The outermost layer, which is solid and relatively thin compared to other layers. It includes both the continental crust (thicker and less dense) and the oceanic crust (thinner and denser). The temperature ranges from about 200°C to 400°C.



- 2. Mantle: Located beneath the crust, the mantle extends to a depth of about 2,900 kilometers. It is composed of semi-solid rock that flows slowly. The temperature increases with depth, ranging from about 500°C near the crust to over 4,000°C near the core-mantle boundary.
- 3. Outer Core: This layer is composed of liquid iron and nickel and extends from a depth of about 2,900 kilometers to 5,150 kilometers. The temperature ranges from about 4,000°C to 6,000°C. The movement of the liquid outer core generates Earth's magnetic field.
- 4. Inner Core: The innermost layer, which is solid and composed primarily of iron and nickel. It extends from a depth of about 5,150 kilometers to the center of the Earth at 6,371 kilometers. The temperature is estimated to be around 5,000°C to 6,000°C, but the immense pressure keeps it in a solid state.

Magnetic field: The Earth's magnetic field is generated by the liquid outer core's flow. The Earth's magnetic field is unstable and can temporarily shut down, causing the north and south poles to reverse.

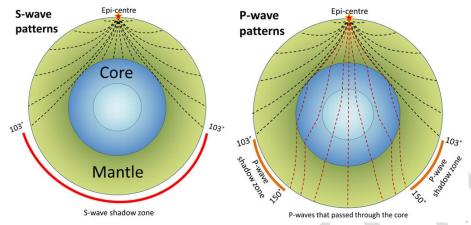
Scientists have learned about the Earth's interior by tracking seismic waves. P-waves slow down at the mantle-core boundary, indicating that the outer core is less rigid than the mantle. S-waves disappear at the mantle-core boundary, indicating that the outer core is liquid.

A seismic shadow zone is an area away from the epicenter where seismic activity is minimized or not present. A seismic shadow zone is not in a constant place; each earthquake



epicenter has a different shadow zone. The finding of shadow zones allowed scientists to prove Earth had an interior layer of liquid.

The seismic shadow zone is the rea of the Earth's surface where seismographs cannot detect an earthquake after the waves have passed through the earth.



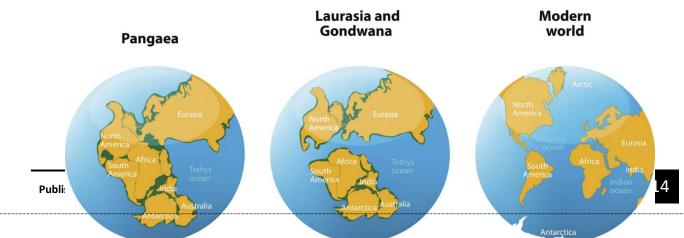
The seismic shadows are the effect of seismic waves striking the core-mantle boundary. P and S waves radiate spherically away from an earthquake's hypocenter (or focus) in all directions and return to the surface by many paths. S waves, however, don't reappear beyond an angular distance of ~103° (as they are stopped by the liquid) and P waves don't arrive between ~103° and 140° due to refraction at the mantle-core boundary.

- The seismic shadow zone is the rea of the Earth's surface where seismographs cannot detect an earthquake after the waves have passed through the earth
- P waves are refracted by the liquid outer core and are not detected between 104° and 140°
- S waves cannot pass through the liquid outer core and are not detected beyond 104°

CONTINENTAL DRIFT

Continental drift is a fascinating geological theory that explains the movement of Earth's continents over geological time.

1. Theory Development: The concept of continental drift was first proposed by Alfred Wegener in 1912. He suggested that the continents were once part of a single supercontinent called Pangaea, which began to break apart around 200 million years ago.





- 2. **Evidence**: Wegener's theory was based on several lines of evidence:
 - Fossil Evidence: Similar fossils of plants and animals were found on continents that are now widely separated by oceans.
 - Geological Evidence: Similar rock formations and mountain ranges were found on continents that are now far apart.
 - Climatic Evidence: Evidence of past climates, such as glacial deposits, indicated that continents were once located in different positions relative to the poles.
- 3. Mechanism: The mechanism of continental drift is the movement of continents across the ocean floor due to the Earth's mantle, which is hot and contains fluids:
 - Plate tectonics: The theory of plate tectonics provided the explanation for continental drift. According to this theory, Earth's lithosphere (the rigid outer layer) is divided into several large and small plates that float on the semi-fluid asthenosphere beneath them. These plates move due to the convective currents in the mantle, causing the continents to drift. The Earth's lithosphere is divided into several plates that float on the asthenosphere, a semi-fluid. The movement of these plates causes the continents to drift.
 - Tectonic forces: The continents can collide or split apart due to tectonic plate movement
 - Earth's rotation: The Earth's rotation causes a centrifugal force that pushes continents away from the poles.
 - Tidal forces: The gravitational pull of the moon and the sun creates tidal forces that can contribute to the movement of continents.
- 4. Modern Understanding: Today, we know that the movement of tectonic plates is responsible for many geological phenomena, including earthquakes, volcanic activity, and the formation of mountain ranges. The study of plate tectonics has revolutionized our understanding of Earth's dynamic nature.

Criticisms of Continental Drift Theory: Wegener discusses how buoyancy, tidal currents, and gravity all play a role. These, however, were insufficient to move continents.

- He favors westward or equatorial travel, but movements have occurred in all directions.
- Alfred Wegener was unable to explain the pre-carboniferous period of time. He didn't say why the drift started in the Mesozoic era and not earlier.
- Oceans were not taken into account in the theory.
- Oceanic ridges and Island arcs were not explained by the idea.
- Large-scale motions are thought to be impossible due to the rigidity of the Earth's crust.
- Wegener's theories failed to provide a convincing mechanism for supporting the displacement of bigger masses during long voyages.

ISOSTASY

Isostasy is the state of gravitational equilibrium between Earth's crust (lithosphere) and mantle such that the crust "floats" at an elevation that depends on its thickness and density. Isostasy is a dynamic process. As the Earth changes, the crust and mantle must adjust to maintain equilibrium.



For example, when glaciers melt, they remove weight from the crust. This causes the crust to rise. Conversely, when glaciers form, they add weight to the crust, causing it to sink. Isostasy can also be affected by tectonic activity, such as the formation of new mountains or the subduction of oceanic crust. The Isostasy term was first used by the American geologist **Dutton (1889).** The theory of plate tectonics (1970s) provides a new framework for understanding Isostasy.

The basic principle of Isostasy is that the lithosphere will adjust its elevation to achieve equilibrium, much like how a floating object in a fluid will adjust its position to balance its weight and buoyancy.

There are two major theories that explain the concept of Isostasy.

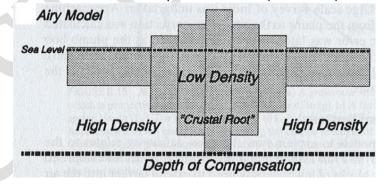
- Airy's views on Isostasy.
- Pratt's views on Isostasy.

Airy's views on Isostasy: The earliest model of isostasy, known as Airy Isostasy, was proposed by George Biddell Airy, a 19th-century British astronomer. The lithosphere, the Earth's outermost shell, is assumed to be a sequence of blocks of constant density in this model. While the density of these pieces remains constant, their thickness changes.

Consider an iceberg drifting in the sea. The "root" of the iceberg lies hidden beneath the water's surface, while the tip protrudes. The deeper the root spreads beneath the surface, the greater the iceberg's tip. Similarly, mountainous regions of the Earth have a thicker part of the crust (or a "root")

stretching down into the denser mantle in Airy Isostasy. This extra "root" serves to balance out the mountain's increased mass above the surface.

When erosion takes down a mountain's mass over time, the crust beneath rises in reaction, maintaining isostatic balance. Airy Isostasy, in essence, defines a 'floating' lithosphere in which the thicker sections



extend deeper into the mantle, just like larger icebergs sink deeper into the sea. This model suggests that the Earth's crust has varying thickness under mountain ranges compared to ocean basins. The thicker portions of the crust (mountains) "float" higher on the asthenosphere, similar to how icebergs with larger submerged parts project more above the water surface.

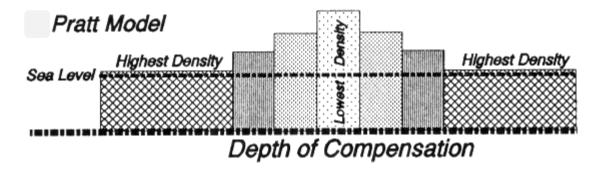
Pratt's views on Isostasy

With his Pratt Isostasy model, British geologist John Henry Pratt took an alternative method to explain isostatic equilibrium. Pratt's hypothesis is based on the density of the lithospheric blocks rather than their thickness.

Consider a similar-sized wooden block and a sponge floating in a tub of water. Despite being the same size as the wooden block, the sponge will float higher since it is less dense. Similarly, in Pratt Isostasy, less dense portions of the Earth's crust (such as those made of less dense rock types or those underlain by substantial sedimentary deposits) 'float' higher on the denser mantle than denser ones.



As a result of these density differences within the crust, Pratt's model implies that the varied elevations we see across the Earth's surface, from plains to plateaus, are the result of these density variations within the crust itself.



The concept of isostasy is central to our understanding of various geologic processes and phenomena, including mountain building, sedimentation, erosion, and sea-level changes. It provides insight into the dynamic and responsive nature of the Earth's crust in relation to the mantle beneath.

- 1. Applications: Isostasy helps explain various geological phenomena, such as the formation of mountain ranges, the subsidence of basins, and the uplift of regions after the melting of ice sheets (a process known as post-glacial rebound).
- 2. **Dynamic Equilibrium:** The Earth's crust is constantly adjusting to maintain isostatic equilibrium. When weight is added or removed (e.g., through erosion, sediment deposition, or glacial melting), the crust responds by sinking or rising to restore balance.

Isostasy is a key concept in understanding the dynamic nature of the Earth's crust and its response to various geological processes.

- Explain the concept of isostasy and discuss its significance in understanding the Earth's crust.
- Compare and contrast the theories of isostasy proposed by Airy and Pratt.
- Discuss the role of isostasy in the formation of various landforms. Provide examples.
- How does isostasy contribute to the understanding of mountain building processes?
- Evaluate the impact of isostatic adjustments on coastal regions.

PLATE TECTONICS

In late 1960's, the consolidation of works and evidences of different geomorphologists like Mc Kenzie, Parker, Wilson, A. Holmes, Herry Hess, Wegner etc. led to the development of plate tectonics theory.

According to Plate tectonics theory, the Earth is made up of 3 main layers (core, mantle, crust). On the surface of the Earth are tectonic plates that slowly move around the globe.

A tectonic plate (lithospheric plate) is a massive, irregularly shaped slab of plate of solid rock generally composed of both continental and oceanic lithosphere.



Under the effect of convective magma current (Convection current Hypothesis by A. Holmes), the lithospheric plate floats over the fluid-like (visco-elastic solid) asthenosphere. That leads to plate movement. The relative movement of the plates typically ranges from zero to 100 mm annually.

The cumulative effects of three phenomena like ridge push, slab pull, frictional drag led to plate movement.

The forces: The forces for plate movement, are generated due to internal heat of Earth that includes

radioactivity, residual heat and gravitational heat.

The earth's lithosphere is composed of seven major plates and many minor plates. The lithosphere is a rigid outermost shell of earth and is broken up into tectonic plates.

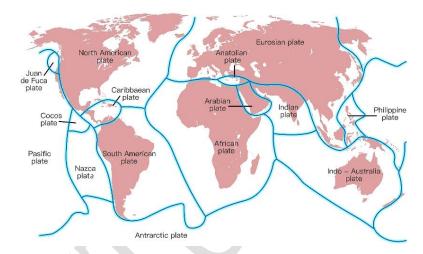


Plate Boundaries: The interactions between tectonic plates occur at plate boundaries, which are classified into three main types:

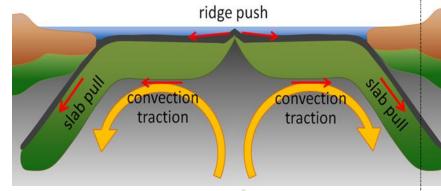
- Divergent Boundaries: Plates move apart from each other, creating new crust as magma rises from the mantle. This process occurs at mid-ocean ridges, such as the Mid-Atlantic Ridge.
- Convergent Boundaries: Plates move towards each other, leading to the subduction of one plate beneath another or the collision of two continental plates.
 This process forms mountain ranges, such as the Himalayas, and volcanic arcs, like the Andes.
- Transform Boundaries: Plates slide past each other horizontally, causing earthquakes along faults, such as the San Andreas Fault in California.
- 2. Mechanism of Plate Movement: The movement of tectonic plates is driven by several forces, including:
- Divergent

 Convergent

 Transform Fault
- Mantle Convection: Heat from the Earth's interior causes convection currents in the mantle, which drag the plates along.
- Ridge Push: The elevated position of mid-ocean ridges creates a gravitational force that pushes plates away from the ridge.



 Slab Pull: The sinking of a denser oceanic plate into the mantle at a subduction zone pulls the rest of the plate along with it.



- 3. Geological Phenomena: Plate tectonics is responsible for many geological phenomena, including:
 - Earthquakes: Caused by the sudden release of energy as plates move past each other.
 - Volcanic Activity: Occurs at divergent and convergent boundaries where magma reaches the surface.
 - Mountain Building: Results from the collision and compression of tectonic plates.
- 4. Historical Development: The theory of plate tectonics was developed in the mid-20th century, building on earlier concepts like continental drift proposed by Alfred Wegener. It revolutionized our understanding of Earth's dynamic processes and provided a unifying framework for geology.

Plate tectonics is a dynamic and ongoing process that shapes the Earth's surface and influences its geological activity

- Explain the theory of plate tectonics and discuss how it has improved upon the continental drift theory.
- Describe the different types of plate boundaries and the landforms associated with each.
- Discuss the role of plate tectonics in the formation of the Himalayas and the Appalachian Mountains.
- How does the theory of plate tectonics explain the occurrence of earthquakes and volcanic activity?
- Evaluate the impact of plate tectonics on the distribution of mineral resources.

MOUNTAIN BUILDING

Mountain building, or orogeny, is a geological process that forms mountain ranges. It involves the deformation of the Earth's crust due to tectonic forces. Here's a detailed overview:

- 1. **Tectonic Plate Interactions**: Mountain building primarily occurs at convergent plate boundaries, where two tectonic plates collide. This collision can involve:
 - Continental-Continental Collision: When two continental plates collide, they
 crumple and fold, creating large mountain ranges like the Himalayas.



 Oceanic-Continental Collision: When an oceanic plate collides with a continental plate, the denser oceanic plate is subducted beneath the continental plate. This process forms volcanic mountain ranges, such as the Andes.

2. Processes Involved:

- Folding: Layers of rock are compressed and folded, creating structures like anticlines (upward folds) and synclines (downward folds).
- Faulting: Rocks break and move along faults, leading to the formation of fault-block mountains. Examples include the Sierra Nevada in the United States.
- Volcanism: Magma from the mantle can rise to the surface, forming volcanic mountains. This process is common at subduction zones and mid-ocean ridges.
- Erosion and Uplift: Once mountains are formed, they are subject to erosion by wind, water, and ice. Erosion wears down the mountains, but isostatic rebound (the rise of the crust after the removal of weight) can cause further uplift, maintaining the elevation of the mountain range.

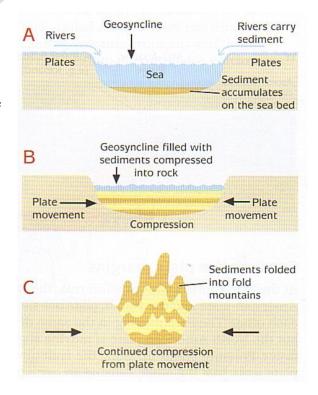
4. Examples of Mountain Ranges:

- o Himalayas: Formed by the collision of the Indian Plate and the Eurasian Plate.
- Andes: Formed by the subduction of the Nazca Plate beneath the South American Plate.
- Rockies: Formed by a combination of subduction and other tectonic processes.

Mountain building is a dynamic and ongoing process that shapes the Earth's surface and creates some of the most dramatic landscapes on our planet.

Geosynclinal Orogen Theory: Leopold Kober, a German geologist, proposed the **Geosynclinal Orogen Theory** to explain mountain building. Here are the key points of his theory:

- Geosynclines: Kober identified mobile zones of water, called geosynclines, where mountains are formed. These zones were surrounded by rigid landmasses known as kratogen.
- 2. **Compression**: According to Kober, the compression of geosynclines situated on the margins of these rigid landmasses leads to the formation of mountain ranges.
- Randketten and Zwischengebirge: The
 mountain ranges formed at the borders of
 these rigid masses are called Randketten. The
 zone between these mountain ranges, which
 remains relatively unaffected by folding, is
 called Zwischengebirge or the Median Mass.





4. **Contraction Hypothesis**: Kober believed that the Earth's cooling and subsequent contraction provided the necessary force for mountain building.

Kober's theory emphasized the role of horizontal movements and compression in the formation of mountains, contrasting with other theories that focused on vertical movements. Today, most geologists consider the geosynclinal theory to be obsolete and instead explain the development of mountain ranges in terms of plate tectonics.

HIMALAYA FORMATION

The Himalayas are bordered to the northwest by the mountain ranges of the Hindu Kush and the Karakoram and to the north by the high and vast Plateau of Tibet. It is a stretch 2,900 km along the border between India and Tibet.

Among the most dramatic and visible creations of plate-tectonic forces are the lofty Himalayas. This immense mountain range began to form between 40 and 50 million years ago, when two large landmasses, India and Eurasia, driven by plate movement, collided. Because both these continental landmasses have about the same rock density, one plate could not be subducted under the other. The pressure of the impinging (colliding) plates could only be relieved by thrusting skyward, contorting the collision zone, and forming the jagged Himalayan peaks.

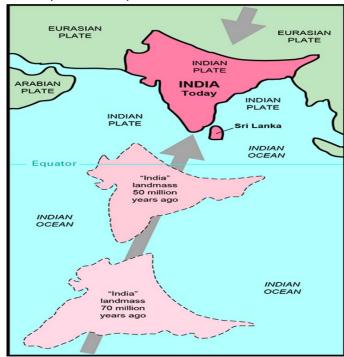
The Indian plate includes **Peninsular India and the Australian continental portions.**The subduction zone along the Himalayas forms the northern plate boundary in the form of **continent convergence.**

Evolution of Himalaya: India was a large island situated off the Australian coast, in a vast ocean. The Tethys Sea separated it from the Asian continent till about 225 million years ago. India is supposed to have started her northward journey about 200 million years ago at the time when Pangaea broke. India collided with Asia about 40-50 million years ago causing rapid uplift of the Himalayas. The positions of India since about 71 million years till the present are shown in the

picture.

It also shows the position of the Indian subcontinent and the Eurasia plate. About 140 million years before the present, the subcontinent was located as south as 50-degree south latitude. The two major plates were separated by the Tethys Sea and the Tibetan block was closer to the Asiatic landmass.

During the movement of the Indian plate towards the Asiatic plate, a major event that occurred was the outpouring of lava and formation of the Deccan Traps. This started somewhere around 60 million years ago and continued for a long period of time. Note that the subcontinent was still close to the equator. From 40 million years ago and





thereafter, the event of formation of the Himalayas took place. Scientists believe that the process is still continuing, and the height of the Himalayas is rising even to this date.

At present, the movement of India continues to put enormous pressure on the Asian continent, and Tibet in turn presses on the landmass to the north that is hemming it in. The net effect of plate-tectonics forces acting on this geologically complicated region is to squeeze parts of Asia eastward toward the Pacific Ocean. One serious consequence of these processes is a deadly "domino" effect: tremendous stresses build up within the Earth's crust, which are relieved periodically by earthquakes along the numerous faults that scar the landscape.

- Explain the process of orogeny and its significance in mountain building.
- Discuss the different types of mountains and the processes involved in their formation.
- Analyze the role of plate tectonics in the formation of the Himalayas.
- How do isostatic adjustments contribute to mountain building?
- Evaluate the impact of mountain building on climate and biodiversity.

VOLCANISM

Volcanism is the process through which magma (molten rock) and gases from the Earth's interior are expelled onto the surface. This process is responsible for the formation of volcanoes and various volcanic features. Here's a detailed overview:

- Magma Formation: Magma forms in the mantle due to the melting of rocks caused by factors such as increased temperature, decreased pressure, or the addition of volatiles (e.g., water).
- 2. Magma Ascent: Once formed, magma rises towards the Earth's surface because it is less dense than the surrounding solid rock. It moves through cracks and fractures in the crust.
- 3. Volcanic Eruptions: When magma reaches the surface, it erupts as lava, along with gases and volcanic ash. Eruptions can vary in intensity, from gentle lava flows to explosive events.

A volcano is an opening in the earth's crust that allows lava, volcanic ash, and gases to escape. There are different types of volcanoes based on the shape, nature of the eruption, frequency of eruption, etc.

Based on the Shape: 4 types of volcanos are found and are discussed below;

Cinder Cones: Cinder cones are round or oval cones made up of tiny lava pieces blown up from a single vent. Examples: Mount Suribachi , Mount Omuro in Japan, Mount Mayabobo in the Philippines

Composite Volcano: Composite volcanoes are steep-sided volcanoes made up of multiple layers of volcanic rocks, most of which are made up of high-viscosity lava, ash, and rock debris. Examples: Mount Saint Helens

Shield Volcano: Shield volcanoes have long, gradual slopes formed by basaltic lava flows and are fashioned like a bowl or shield in the middle. Examples: Mauna Loa and Mauna Kea.



Lava Domes: Lava domes arise when erupting lava becomes too thick to flow and stacks up near the volcanic vent, forming a steep-sided mound. Examples: Santa María lava dome (Guatemala), Sollipulli lava dome (Chile)

Based on the Nature of the Eruption: 5 types of volcanos are found and are discussed below;



- Hawaiian Volcanoes: Hawaiian Volcanoes are named after the Kilauea Volcano on Hawaii's Big Island which is known for its stunning fire fountains.
- Strombolian Eruption: Strombolian eruptions are fluid lava bursts from the mouth of a magma-filled summit conduit.
- Vulcanian Eruption: A Vulcanian eruption is a viscous magma explosion that is short, intense, and relatively tiny.
- Plinian Eruptions: Plinian eruptions are the largest and most violent of all volcanic eruptions.
- Vesuvian Volcanoes: Vesuvian Volcanoes have magma that is ejected from a start cone vent and are particularly violent and explosive.

Based on the Frequency of Eruption

- Active Volcano: Active volcanoes are volcanoes that are either erupting or on the verge of
 eruption. Mount Etna (Italy), Hawaiin Islands (Pacific Ocean), Mauna Loa (Pacific Ocean),
 Mount Vesuvius (Italy), and Barren Island (India) are some examples of Active Volcanoes
 around the world.
- Dormant Volcano: A dormant volcano is one that is not erupting at the present but has erupted in the past and is expected to erupt again. Example: Mauna Kea, Mt. St. Helens. Narcondam Island (Andaman Sea)
- Extinct Volcano: Extinct volcanoes are considered to be dormant and unlikely to erupt again. Example: Kohala, the Big Island of Hawaii's oldest volcano, Aconcagua of the Andes is a typical example of an extinct volcano.

4. Volcanic Features:

- o Lava Flows: Streams of molten rock that flow from a volcanic vent.
- Pyroclastic Flows: Fast-moving currents of hot gas and volcanic matter that can be extremely destructive.



- Volcanic Ash: Fine particles of volcanic rock and glass that are ejected during an eruption and can travel long distances.
- Calderas: Large, basin-like depressions formed when a volcano collapses after a massive eruption.
- 5. Volcanic Hazards: Volcanic eruptions can pose significant hazards, including lava flows, ashfall, pyroclastic flows, and volcanic gases. These hazards can impact human populations, infrastructure, and the environment.
- 6. Volcanic Benefits: Despite the hazards, volcanism also has benefits. Volcanic soils are often very fertile, supporting agriculture. Volcanic activity can also create geothermal energy resources and contribute to the formation of valuable mineral deposits.

Volcanism is a dynamic and powerful process that shapes the Earth's surface and has both destructive and beneficial impacts.

EARTHQUAKES

Earthquakes are natural phenomena that occur when there is a sudden release of energy in the Earth's crust, causing the ground to shake. Here's a detailed overview:

- Causes: Earthquakes are primarily caused by the movement of tectonic plates. The Earth's
 crust is divided into several plates that float on the semi-fluid mantle beneath them. When
 these plates interact at their boundaries, stress builds up until it is released as seismic
 energy. Other causes include volcanic activity, landslides, and human activities like mining
 and reservoir-induced seismicity.
- 2. Focus and Epicenter: The point within the Earth where the earthquake originates is called the focus or hypocenter. The point directly above it on the surface is the epicenter. The energy released from the focus travels in the form of seismic waves.
- 3. Seismic Waves: There are different types of seismic waves:
 - Primary (P) Waves: These are the fastest seismic waves and travel through solids, liquids, and gases. They compress and expand the material they move through.
 - Secondary (S) Waves: These are slower than P waves and can only travel through solids. They move the ground up and down or side to side.
 - Surface Waves: These travel along the Earth's surface and are responsible for most
 of the damage during an earthquake. They include Love waves and Rayleigh waves.
- 4. Measurement: Earthquakes are measured using seismographs, which record the intensity and duration of seismic waves. The magnitude of an earthquake is commonly measured on the Richter scale or the moment magnitude scale (Mw). The intensity, which measures the effects of an earthquake, is often assessed using the Modified Mercalli Intensity (MMI) scale.
- 5. Effects: Earthquakes can cause a wide range of effects, including ground shaking, surface rupture, landslides, tsunamis, and liquefaction. The severity of these effects depends on the earthquake's magnitude, depth, distance from the epicenter, and local geological conditions.



 Preparedness and Mitigation: To reduce the impact of earthquakes, it is important to have proper building codes, early warning systems, and emergency preparedness plans.
 Education and awareness about earthquake safety can also help communities respond effectively during an event.

Earthquakes are a powerful reminder of the dynamic nature of our planet.

WEATHERING AND EROSION

Weathering and erosion are two fundamental geological processes that shape the Earth's surface. Here's a detailed overview:

- Weathering: Weathering is the process of breaking down rocks and minerals into smaller pieces through physical, chemical, and biological means. It occurs in place, without the movement of materials.
 - Physical Weathering: Also known as mechanical weathering, this process involves the physical breakdown of rocks without changing their chemical composition.
 Examples include:
 - Frost Wedging: Water seeps into cracks in rocks, freezes, and expands, causing the rocks to break apart.
 - Thermal Expansion: Repeated heating and cooling of rocks cause them to expand and contract, leading to fractures.
 - Abrasion: Rocks and sediments grind against each other, wearing down surfaces.
 - Chemical Weathering: This process involves the chemical alteration of minerals within rocks, leading to their breakdown. Examples include:
 - Oxidation: The reaction of minerals with oxygen, forming oxides (e.g., rusting of iron).
 - Hydrolysis: The reaction of minerals with water, forming new minerals and soluble ions.
 - Carbonation: The reaction of minerals with carbonic acid (formed from carbon dioxide and water), leading to the dissolution of minerals like limestone.
 - Biological Weathering: This process involves the breakdown of rocks by living organisms. Examples include:
 - Root Wedging: Plant roots grow into cracks in rocks, exerting pressure and causing them to break apart.
 - Lichen and Moss: These organisms produce acids that chemically weather rocks
- 2. **Erosion**: Erosion is the process of transporting weathered materials from one location to another by natural agents such as water, wind, ice, and gravity.



- Water Erosion: Running water, such as rivers and streams, can carry away sediments and shape landscapes through processes like:
 - Rill and Gully Erosion: Small channels (rills) and larger channels (gullies) are formed by flowing water.
 - River Erosion: Rivers carve valleys and canyons, transport sediments, and deposit them in deltas and floodplains.
- Wind Erosion: Wind can transport fine particles like sand and silt, leading to the formation of features such as:
 - Dunes: Mounds of sand formed by wind deposition.
 - Deflation: The removal of loose particles from the ground, leaving behind a surface of larger, more resistant rocks.
- o **Ice Erosion**: Glaciers can erode landscapes through processes like:
 - Plucking: Glaciers pick up and transport rocks and sediments.
 - Abrasion: Rocks embedded in glaciers grind against the bedrock, carving out valleys and fjords.
- Gravity Erosion: Gravity can cause the movement of materials downslope through processes like:
 - Landslides: Rapid movement of rock and soil down a slope.
 - Creep: Slow, gradual movement of soil and rock down a slope.

Weathering and erosion work together to shape the Earth's surface, creating diverse landscapes and contributing to the rock cycle.

GEOMORPHIC CYCLES

Geomorphic cycles, also known as the cycle of erosion, describe the stages of landscape evolution and the processes that shape the Earth's surface over geological time.

DAVIS GEOMORPHIC CYCLES

The concept was first introduced by William Morris Davis in the late 19th century. Davis has described the landscape as the function of structure, process and stage. This means that in a landscape all these three-play dominant role.

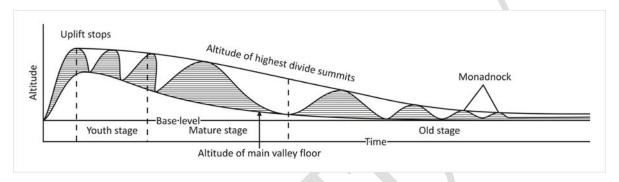
- Structure: Which includes 'nature' (hardness, permeability) and 'attitude' (folds, faults, joints, slopes) of rocks?
- Process: Implies the factors or agents responsible for weathering and erosion.
- Time: Implies the stage at which the cycle is—youth, maturity or old age.

There are several **assumptions** that underlie the Davis Cycle of Erosion:



- The landforms evolved through the interaction of Endogenetic and exogenetic forces.
- The development of landforms takes place in an orderly manner in response to environmental changes.
- Streams erode the valley downwards till the graded condition is achieved.
- The upliftment of landmass is a very rapid process.
- Erosion does not start till the upliftment gets completed.
- The cycle assumes that the underlying geology and climate of a region are stable over the timescale of the cycle.

His model describes the stages of landscape evolution and the processes that shape the Earth's surface over geological time. Here's a detailed overview of Davis's concepts:



1. Youthful Stage:

- Characteristics: In this initial stage, the landscape is characterized by steep slopes, deep valleys, and rapid river flow. The terrain is rugged, with little to no floodplains.
- Processes: Dominated by vertical erosion, rivers cut deep valleys and create Vshaped profiles. Waterfalls and rapids are common.

2. Mature Stage:

- Characteristics: The landscape becomes more subdued as the slopes become gentler and valleys widen. River meanders and floodplains start to develop.
- o **Processes:** Lateral erosion becomes more prominent, widening the valleys and creating meanders. Sediment deposition begins to play a significant role.

3. Old Stage:

- Characteristics: The landscape is characterized by broad, flat valleys, extensive floodplains, and meandering rivers. The terrain is relatively flat with low relief.
- Processes: Erosion and deposition reach a balance, with rivers depositing sediments to form features like oxbow lakes and natural levees.

4. Rejuvenation:

- Characteristics: If there is a change in base level (e.g., due to tectonic uplift or sealevel change), the landscape can be rejuvenated, leading to renewed erosion.
- Processes: Rivers may incise into their floodplains, creating terraces and new valley profiles.



5. **Dynamic Equilibrium**:

- Characteristics: Modern geomorphology recognizes that landscapes are in a state of dynamic equilibrium, constantly adjusting to changes in climate, tectonics, and other factors.
- Processes: Erosion, transportation, and deposition processes continuously shape the landscape, maintaining a balance over time.

Davis's model provides a framework for understanding the long-term evolution of landscapes and the processes that drive these changes.

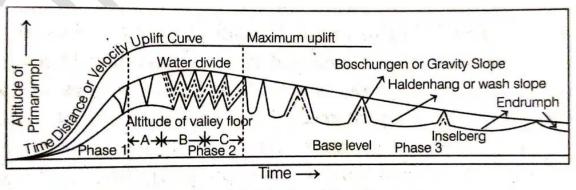
PENCK'S GEOMORPHIC CYCLES

Walther Penck, a German geomorphologist, proposed an alternative model to William Morris Davis's cycle of erosion in 1924. Penck's model focuses on the relationship between tectonic activity and erosion, emphasizing the continuous interaction between uplift and degradation. Here are the key concepts of Penck's geomorphic cycles:

Continuous Uplift and Erosion: Unlike Davis's model, which is time-dependent and follows
distinct stages (youth, maturity, and old age), Penck's model suggests that landform
development is a continuous process. The landscape evolves through the interplay of uplift
(tectonic activity) and erosion (exogenetic processes).

2. Phases of Development:

- Aufsteigende Entwicklung (Waxing or Accelerated Rate of Development): This
 phase occurs when the rate of uplift exceeds the rate of erosion, leading to the
 formation of steep slopes and high relief.
- Gleichformige Entwicklung (Uniform Rate of Development): In this phase, the rates
 of uplift and erosion are balanced, resulting in a stable landscape with moderate
 slopes.
- Absteigende Entwicklung (Waning or Decelerating Rate of Development): This
 phase occurs when the rate of erosion exceeds the rate of uplift, leading to the
 reduction of relief and the formation of gentler slopes.



Penck's Cycle of Erosion



- 3. Primarumpf: Penck introduced the concept of "primarumpf," which refers to the initial surface or primary peneplain before uplift. This surface represents a featureless landmass that undergoes uplift and subsequent erosion.
- 4. **Dynamic Equilibrium**: Penck's model emphasizes that landscapes are in a state of dynamic equilibrium, constantly adjusting to changes in tectonic activity and erosion rates. This continuous interaction shapes the landforms over time.

Penck's model provides a more flexible and dynamic understanding of landscape evolution compared to Davis's cyclic model.

Positive Points:

- 1. Penck followed a deductive approach and did not restrict himself to any particular condition.
- 2. Compared to the Davisian cycle, Penck's approach was forward looking.
- 3. Penck, quite appropriately, emphasised the mutual relation between uplift and the deepening of valleys. This indicates Penck's respect for geological evidence. Penck's third stage is evident in the Middle Alps.

Drawbacks:

- 1. Penck gave too much importance to the role of endogenetic forces.
- 2. The orderliness in landform changes, as assumed by Penck, may be difficult to achieve.
- 3. Inadequate knowledge about the initial pristine landscape does not allow much verification.
- 4. The concept of geographical cycle of erosion itself has been criticised by many, since many of the cyclic generalisations are based on untested assumptions. An overemphasis on historical and evolutionary studies in landforms results in the reconstruction of stages of evolution becoming the focus of study.

FLUVIAL LANDFORMS

Fluvial landforms are created by the action of rivers and streams. These landforms are shaped by processes such as **erosion**, **transportation**, and **deposition of sediments**.

Fluvial landforms are landforms created by rivers and streams, and are generally categorized as either erosional or depositional:

Erosional landforms: Created when rivers wear away the land through erosion:

- Valleys
- Gulleys
- Entrenched meanders
- River terraces
- Plungepools
- Gorges
- Canyons
- Waterfalls

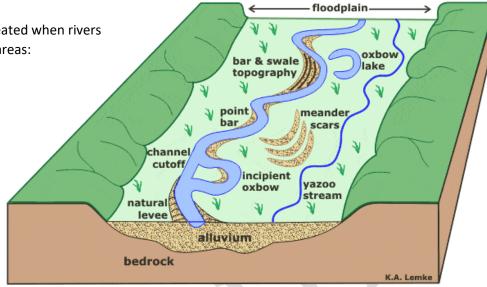


Rapids

Depositional landforms: Created when rivers deposit sediments on plain areas:

- Deltas
- Alluvial fans
- Flood plains
- Meanders
- Oxbow lakes
- Natural levees
- Braided channels

Here are some key fluvial landforms:



1. River Valleys:

These are elongated depressions carved by rivers. They can be V-shaped in youthful stages or U-shaped in mature stages.

- 2. Floodplains: Flat areas adjacent to rivers that are periodically flooded. They are formed by the deposition of sediments during floods.
- 3. Meanders: Curved bends in a river. They form as the river erodes the outer banks and deposits sediments on the inner banks.
- 4. Oxbow Lakes: Crescent-shaped lakes formed when a meander is cut off from the main river channel.
- 5. Levees: Raised banks along a river, formed by the deposition of sediments during floods.
- 6. **Deltas**: Landforms created at the mouth of a river where it deposits sediments as it enters a standing body of water, like an ocean or a lake.
- 7. Alluvial Fans: Fan-shaped deposits of sediments formed where a river flows out of a mountainous area onto a flatter plain.
- 8. Terraces: Step-like landforms along the sides of a river valley, formed by the river cutting down into its floodplain.

These landforms illustrate the dynamic nature of rivers and their ability to shape the landscape over time.

AEOLIAN LANDFORMS

Aeolian landforms are landforms created by wind action in arid, semi-arid, and extremely dry regions. Wind is the primary geomorphic force in these regions, and the landforms it creates are the result of its erosional and depositional activities.

Aeolian landforms are created by the erosional and depositional actions of wind, and include a variety of features:

Erosional features: These include:

- Yardangs
- Zuegen



- Desert pavements
- **Deflation hollows**
- Ventifacts
- Mesas
- **Buttes**
- Mushroom rocks

Depositional features: These include:

- Sand dunes
- Loess deposits
- Sand ripples
- Sand ridges
- Sandfields
- Sand shadows
- Sand drifts

Aeolian Landforms: Erosional



Zeugen















Blow Outs

Inselbergs

Desert Pavement

Longitudinal Dunes Transverse Dunes

Star Dunes

- 1. Desert Pavement: A surface covered with closely packed, interlocking angular or rounded rock fragments. It forms as finer particles are removed by wind and water, leaving behind a layer of larger rocks.
- 2. Sand Dunes: Mounds or ridges of sand formed by wind deposition. They come in various shapes, including:
 - Barchan Dunes: Crescent-shaped dunes with horns pointing downwind.
 - Transverse Dunes: Long ridges perpendicular to the prevailing wind direction.
 - Star Dunes: Radial dunes with multiple arms extending from a central point.
- Yardangs: Streamlined, wind-eroded ridges that are aligned with the prevailing wind direction. They are formed by the abrasion of softer rock layers, leaving behind more resistant ridges.
- 4. Playas: Flat, dry lake beds found in desert basins. They form when temporary lakes evaporate, leaving behind a layer of salt, clay, or silt.
- 5. Alluvial Fans: Fan-shaped deposits of sediments formed where a river flows out of a mountainous area onto a flatter plain. They are common in arid regions where intermittent streams deposit sediments.
- 6. Wadis: Dry riverbeds or valleys that only contain water during periods of heavy rain. They are common in arid and semi-arid regions.
- 7. Mesas and Buttes: Flat-topped, steep-sided landforms that are remnants of larger plateaus. Mesas are larger, while buttes are smaller and more isolated.



8. Badlands: Eroded landscapes with steep slopes, minimal vegetation, and intricate networks of gullies and ravines. They form in areas with soft sedimentary rocks and sparse vegetation.

These landforms illustrate the unique processes and features of arid environments.

GLACIAL LANDFORMS

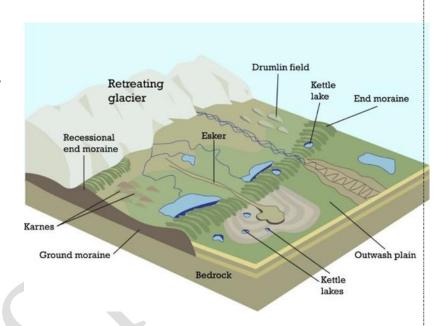
Glacial landforms are created by the movement and melting of glaciers, which shape the landscape through processes of erosion, transportation, and deposition.

Erosional landforms due to Glaciers

- 1. Cirque or Corris
- **2.** Hanging Valleys or U-shaped Valleys, Fjords/fiords
- 3. Horns and Aretes

Depositional Landforms due to Glaciers

- 1. Moraines
- 2. Eskers
- 3. Drumlins
- 4. Kettle hole:



Here are some key glacial landforms:

- 1. Cirques: Bowl-shaped depressions found at the head of a glacier. They are formed by the erosive action of ice and freeze-thaw processes.
- 2. Arêtes: Sharp ridges that form between adjacent cirques or glacial valleys. They are created by the erosion of two glaciers on either side of a ridge.
- 3. Horns: Pyramid-shaped peaks that form when several cirques erode a mountain from different sides. The Matterhorn in the Alps is a famous example.
- U-Shaped Valleys: Also known as glacial troughs, these valleys have a characteristic U-shape, formed by the erosive action of a glacier moving through a pre-existing V-shaped river valley.
- 5. Hanging Valleys: Smaller valleys that join a main glacial valley at a higher elevation. They often form waterfalls where the smaller valley meets the main valley.
- 6. Moraines: Accumulations of rock and debris deposited by a glacier. They can be found at the sides (lateral moraines), at the end (terminal moraines), or beneath (ground moraines) the glacier.



- 7. **Drumlins**: Streamlined, elongated hills composed of glacial till. They are formed beneath a glacier and indicate the direction of ice movement.
- 8. Eskers: Long, winding ridges of sand and gravel deposited by meltwater streams flowing beneath a glacier.
- 9. **Kettles**: Depressions formed by the melting of large blocks of ice left behind by a retreating glacier. They often fill with water to form kettle lakes.
- 10. Outwash Plains: Flat areas of sediment deposited by meltwater streams in front of a glacier. They are composed of sorted materials like sand and gravel.

These landforms illustrate the powerful impact of glaciers on shaping the Earth's surface.

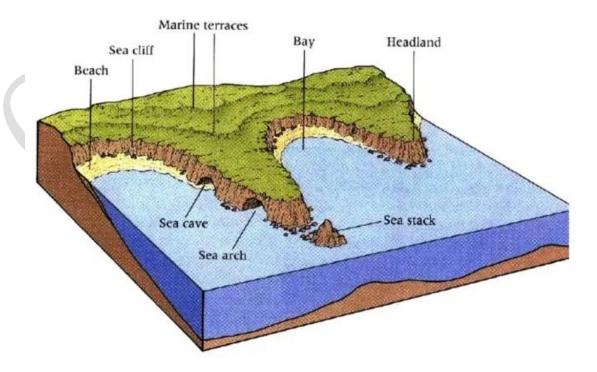
COASTAL LANDFORMS

Coastal landforms are the physical features of a coast, which are created by a combination of processes, sediments, and the geology of the coast. Coastal landforms are shaped by the interaction of the ocean and the land, influenced by processes such as wave action, tides, and currents. They can be formed by a variety of processes, including:

Deposition: When a body of water loses energy, sediments like sand and silt settle on a surface, creating new landforms. Deposition can occur when waves enter a shallow area, hit a sheltered area, or when there is a weak wind.

Erosion: When waves crash against rock faces, the movement of rock creates coastal erosion landforms like cliffs and wave cut platforms.

Weathering: Over time, headlands, which are areas of harder rock, are slowly eroded by wind, rain, and waves. This process can create caves, arches, stacks, and bays.





- Beaches: Sandy or pebbly shores formed by the deposition of sediments carried by waves and currents. They can vary in size and shape depending on the wave energy and sediment supply.
- 2. Cliffs: Steep rock faces along the coast, formed by the erosive action of waves. Over time, waves erode the base of the cliff, causing it to retreat inland.
- Headlands and Bays: Headlands are protruding landforms that extend into the sea, while bays are recessed areas between headlands. Headlands are typically composed of more resistant rock, while bays are formed in areas of softer rock.
- Sea Arches and Sea Stacks: Sea arches are natural rock formations created by the
 erosion of headlands. When the arch collapses, it leaves behind isolated columns of
 rock called sea stacks.
- 5. Spits: Narrow, elongated landforms that extend from the coast into the sea, formed by the deposition of sediments carried by longshore drift. They often have a curved end due to changes in wave direction.
- 6. Barrier Islands: Long, narrow islands parallel to the coast, formed by the accumulation of sand and sediments. They act as a barrier, protecting the mainland from wave action and storms.
- 7. **Estuaries**: Coastal inlets where freshwater from rivers mixes with saltwater from the sea. They are often characterized by rich biodiversity and serve as important habitats for various species.
- 8. Lagoons: Shallow bodies of water separated from the sea by a barrier such as a sandbar or coral reef. They are often found behind barrier islands or within atolls.
- 9. Tidal Flats: Flat, muddy areas exposed during low tide and submerged during high tide. They are typically found in sheltered coastal areas such as estuaries and bays.

These landforms illustrate the dynamic nature of coastal environments and the processes that shape them.

KARST LANDFORMS

Karst topography refers to natural landscape that is largely the result of chemical weathering by water, resulting in caves, sinkholes, cliffs, and steep-sided hills called towers. These features form when water picks up carbon dioxide from the atmosphere and ground to form carbonic acid.

Karst landforms are created by the dissolution of soluble rocks, such as limestone, dolomite, and gypsum, by natural waters. These processes result in unique and often dramatic landscapes. Here are some key karst landforms:

 Sinkholes: Depressions or holes in the ground caused by the collapse of a surface layer. They can vary in size from a few meters to several hundred meters in diameter and depth.



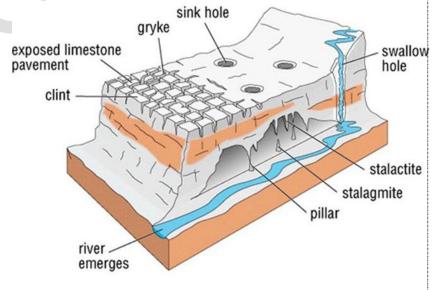
- 2. Caves: Natural underground spaces formed by the dissolution of rock. They can be extensive and complex, with features such as stalactites (hanging from the ceiling) and stalagmites (rising from the floor).
- 3. **Karst Towers**: Tall, steep-sided hills formed by the dissolution of limestone in tropical regions. They are often isolated and can create striking landscapes.
- 4. **Disappearing Streams**: Streams that vanish into the ground, flowing into underground passages and re-emerging elsewhere.
- 5. **Springs**: Natural outlets where groundwater flows to the surface. In karst regions, springs can be large and discharge significant amounts of water.
- 6. Limestone Pavements: Flat, exposed surfaces of limestone that have been eroded to create a pattern of fissures (grikes) and blocks (clints).
- 7. **Dolines**: Large, shallow depressions formed by the dissolution of limestone. They are similar to sinkholes but generally larger and more gently sloping.
- 8. Poljes: Large, flat-floored depressions in karst regions, often with a seasonal or permanent lake. They are formed by the collapse of large cave systems or the dissolution of limestone over a wide area.
- 9. **Karst Valleys**: Valleys formed by the dissolution of limestone, often with steep sides and a flat floor. They can be dry or contain intermittent streams.
- 10. **Stalactite:** A portion of the roof hangs on the roof and on evaporation of water, a small deposit of limestone is left behind contributing to the formation of a stalactite, growing downwards from the roof.

11. **Stalagmite**: The remaining portion of the drop falls to the floor. This also evaporates, leaving behind a small deposit of limestone aiding the formation of a stalagmite,

thicker and flatter, rising upwards from the floor.

These landforms illustrate the unique processes and features of karst landscapes.

The degree of development of karst landforms varies greatly from region to region. Large drainage systems in karst areas are likely to have both fluvial (surface) and karst (underground) drainage components.



Sequential development of landforms in Karst region:

 Conditions that promote karst development are well-jointed, dense limestone near the surface; moderate to heavy rainfall; and good groundwater circulation.



- Limestone (calcium carbonate) dissolves relatively easily in slightly acidic water, which occurs widely in nature.
- Rainwater percolates along both horizontal and vertical cracks, dissolving the limestone and carrying it away in solution.
- Limestone pavements are produced by the removal of surface material, and the vertical fissures along joints are gradually widened and deepened, producing a grooved and jagged terrain.
- As it flows along cracks underground, the water continues to widen and deepen the cracks until they become cave systems or underground stream channels into which narrow vertical shafts may open.

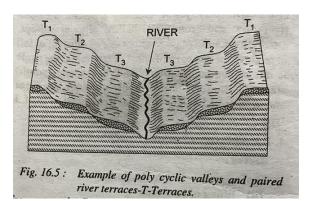
POLYCYCLIC LANDFORMS

A landform that is created by a number of geographical cycles occurring one after the other in order, leaving their distinct marks on the landscape, is called a polycyclic landform. Processes (weathering, erosion and sediment deposition etc.) acting on landforms can also change over time, and a single landscape can undergo several cycles of development. This type of landscape development is polycyclic.

There are various reasons for the formations of these reliefs.

- **Dynamic reasons** involving uplift or subsidence of land resulting in a change in base level. Such changes are mostly localised.
- Eustatic reasons implying a worldwide change in sea level due to diastrophism or glaciation.
- **Static reasons**, e.g. a reduction in river load or an increase in volume (due to precipitation or deforestation) may alter the rate of erosion.
- Climatic reasons, such as aridity, glaciation etc.

These typical landforms resulting from interruptions in the fluvial cycle of erosion and from rejuvenation resulting in the formation of mossaic of poly or multi-cyclic landforms include topographic discordance, valley in valley or multi-storeyed valleys, up-lifted peneplains, incised meanders, paired terraces, nick points etc.



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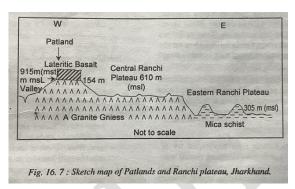
Topographical Discordance : Topographic discordance refers to the creation of older topographic forms above and younger forms below. Topographical discordance leads to formation of **Valley in Valley topography' or paired terraces.** Basically this happens due to



rise in the land of coarse of river. Paired terraces are also significant features of poly-cyclic reliefs. Three pairs of terraces are found in the Himalayas indicating three phases of uplift and consequent rejuvenation.

Example- Almost all Uttarakhand rivers have 3 paired terraces reflecting 3 phases of upliftment. The Damodar valley two-storeyed valley.

Uplifted Peneplain: Uplifted peneplains are formed due to interruption caused by rejuvenation consequent upon regional upliftment. The uplifted peneplains are represented by their remnants of accordant summit levels which rise above the general ground surface of the present-day planation surface. Uplifted peneplains are in fact the results of successive cycles of



erosion wherein several fluvial cycles of erosion are completed in succession.

Example: - The Patlands of the Ranchi plateau is a typical example of uplifted peneplain which is higher than the central Ranchi plateau.

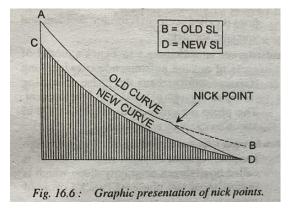
Incised Meanders: Incised meanders are the representative features of rejuvenation and polycyclic reliefs and are developed through vertical erosion leading to valley incision consequent upon renewed erosive power due to rejuvenation.

Example: The Damodar gorge near Rajroppa and Bheraghat gorge of the Narmada near Jabalpur(M.P.)



Knick Point: This is one of the most important feature of Polycyclic land-forms. It represents breaks in slope in the longitudinal profile of river. This is formed due to rejuvenation which lead to fall in the longitudinal of river hence, this is also called as 'head of rejuvenation' which registers gradual recession. Nick Point leads to formation of waterfalls which is also known as 'Nick Point falls' or 'Nick falls'

Example: Hundru falls (76.67 m) on the Subarnarekha river , Dhunwadhar falls on the



- 1. Discuss with examples the influence of vulcanism and diastrophism on the evolution of landscape.
- 2. Define the concept of isostasy and discuss the postulations of Airy and Pratt.
- 3. Critically evaluate the continental drift hypothesis of A. Wegener.
- 4. Role of seismic waves in the study of earth's interior.
- 5. "Structure is a dominant control factor in the evolution of landforms."



- Discuss with suitable examples.
- 6. Describe the landforms which are products of endogenetic forces.
- 7. Explain the sequential development of landforms associated with the coastal areas.
- 8. Provide a critique of the 'geographical cycle' model propounded by Davis.
- 9. With reference to the theory of Plate Tectonics, explain the origin and growth of the Young Fold Mountain Systems of the world.
- 10. Bring out the distinctions between the continental drift theory and the plate tectonics theory.
- 11. Discuss the concept of volcanicity and show how the theory of plate tectonics explains the mechanism of volcanism and volcanic eruption.
- 12. Discuss the process of mechanical and chemical weathering and show their relationship with soil formation.
- 13. Discuss the limitations of the theory of Continental Drift and show how the theory of Plate Tectonics is an improvement over it.
- 14. Discuss the role of Slope, Altitude and Relief (SAR) in landscape development.
- 15. With the help of suitable sketches describe the mountain genesis and mountain types. Give suitable examples from various mountain systems of the world.
- 16. The concept of Plate tectonics has been derived from the isostasy and continental drift. Elaborate citing suitable examples.
- 17. Write an essay on the evolution of continents and oceans using various theories and models.
- 18. Discuss the problems of erosional surfaces and explain the different methods to identify them with suitable diagrams.
- 19. Compare and contrast different types of plate boundaries.
- 20. Discuss the methods of measuring the intensity and magnitude of earthquakes. How are seismic zones demarcated?
- 21. "Landscape is a function of structure, process and stage." Critique the statement.
- 22. "Geomorphological changes are largely responsible for environmental hazards in the Himalayan region." Comment with relevant examples.
- 23. "The knowledge of slope analysis has limited field application in the slope management." Explain.
- 24. Explain the concept of dynamic equilibrium in geomorphology.
- 25. Discuss the role of climate in the development of landforms.
- 26. Describe the various types of mass movements and their impact on the landscape.
- 27. Explain the concept of peneplain and its significance in geomorphology.
- 28. Discuss the role of rivers in shaping the landscape.
- 29. Explain the process of karst topography formation.
- 30. Describe the different types of glaciers and their impact on the landscape.
- 31. Discuss the role of wind in the formation of desert landforms.
- 32. Explain the concept of base level and its significance in geomorphology.
- 33. Discuss the role of tectonic activity in the formation of landforms.
- 34. Explain the process of soil formation and its significance in geomorphology.



(2011)

- 35. Describe the different types of volcanic landforms and their impact on the landscape.
- 36. Discuss the role of human activities in the modification of landforms.
- 37. Explain the concept of geomorphic cycle and its significance in geomorphology. (2009)
- 38. Describe the different types of coastal landforms and their impact on the landscape.
- 39. Discuss the role of weathering in the formation of landforms
- 40. Explain the process of erosion and its significance in geomorphology.